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Data Access Statement: The vegetation census and plant functional traits data that support the findings of this study are available from (28), www.ForestPlots.net (29), and their other original sources. Given data sovereignty from the original data owners, raw data on vegetation censuses across time are not publicly available but can be requested by contacting all researchers through the ForestPlots (30) data request protocol described in forestplots.net/en/join-forestplots/working-with-data. Raw climate data can be accessed through the TerraClimate database at <https://www.climatologylab.org/terraclimate.html> (48). The SPEI data can be obtained from the SPEI database at <https://spei.csic.es/> (83). The computer code used to reproduce the main findings in this manuscript (84) and the plot-level processed data (85) are archived in the Zenodo repository at <https://zenodo.org/>.

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Tropical forests in the Americas are changing too slowly to track climate change

Authors: Jesús Aguirre-Gutiérrez^{1, 120, *}, Sandra Diaz^{2,3}, Sami Rifai⁴, Jose Javier Corral-Rivas⁵, Maria Guadalupe Nava-Miranda^{6,7}, Roy González-M^{8, 119}, Ana Belén Hurtado-M⁸, Norma Salinas Revilla⁹, Emilio Vilanova¹⁰, Everton Almeida¹¹, Edmar Almeida de Oliveira¹², Esteban Alvarez-Davila¹³, Luciana F. Alves¹⁴, Ana Cristina Segalin de Andrade¹⁵, Antonio Carlos Lola da Costa¹⁶, Simone Aparecida Vieira¹⁷, Luiz Aragão^{18,19}, Eric Arets²⁰, Gerardo A. Aymard C.²¹, Fabrício Baccaro²², Yvonne Vanessa Bakker²³, Timothy Russell Baker²⁴, Olaf Bánki²⁵, Christopher Baraloto²⁶, Plínio Barbosa de Camargo²⁷, Erika Berenguer^{1, 28}, Lilian Blanc^{29, 30}, Damien Bonal³¹, Frans Bongers³², Kauane Maiara Bordin³³, Roel Brienen²⁴, Foster Brown³⁴, Nayane Cristina C. S. Prestes¹², Carolina V. Castilho³⁵, Sabina Cerruto Ribeiro³⁶, Fernanda Coelho de Souza³⁷, James A. Comiskey^{38,39}, Fernando Cornejo Valverde⁴⁰, Sandra Cristina Müller³³, Richarly da Costa Silva⁴¹, Julio Daniel do Vale⁴², Vitor de Andrade Kamimura^{23, 43}, Ricardo de Oliveira Perdiz^{44, 45}, Jhon del Aguila Pasquel^{46, 47}, Géraldine Derroire⁴⁸, Anthony Di Fiore^{49, 50}, Mathias Disney^{51, 52}, William Farfan-Rios^{53, 54}, Sophie Fauset⁵⁵, Ted Feldpausch⁵⁶, Rafael Flora Ramos²³, Gerardo Flores Llampazo⁴⁶, Valéria Forni Martins^{57, 58}, Claire Fortunel⁵⁹, Karina Garcia Cabrera⁶⁰, Jorcely Gonçalves Barroso⁶¹, Bruno Héroult^{29, 62}, Rafael Herrera⁶³, Eurídice Nora Honorio Coronado⁶⁴, Isau Huamantupa-Chuquimaco^{65, 66}, John J. Pipoly^{67, 68}, Katia Janaina Zanini³³, Eliana Jiménez⁶⁹, Carlos A. Joly⁵⁷, Michelle Kalamandeen⁷⁰, Joice Klipel³³, Aurora Levesley²⁴, Wilmar Lopez Oviedo^{71, 72}, William E. Magnusson⁷³, Rubens Manoel dos Santos⁷⁴, Beatriz Schwantes Marimon¹², Ben Hur Marimon-Junior¹², Simone Matias de Almeida Reis^{12, 36}, Omar Aurelio Melo Cruz⁷⁵, Abel Monteagudo Mendoza^{54, 106}, Paulo Morandi¹², Robert Muscarella⁷⁶, Henrique Nascimento⁷⁷, David A. Neill⁷⁸, Imma Oliveras Menor^{1, 59}, Walter A. Palacios⁷⁹, Sonia Palacios-Ramos⁸⁰, Nadir Carolina Pallqui Camacho^{24, 81}, Guido Pardo⁸², R. Toby Pennington^{83, 84}, Luciana de Oliveira Pereira⁵⁶, Georgia Pickavance²⁴, Rayana Caroline Picolotto³³, Nigel C. A. Pitman⁸⁵, Adriana Prieto⁸⁶, Carlos Quesada⁸⁷, Hirma Ramírez-Angulo⁸⁸, Maxime Réjou-Méchain⁸⁹, Zorayda Restrepo Correa⁹⁰, José Manuel Reyna Huaymacari⁴⁶, Carlos Reynel Rodriguez⁹¹, Gonzalo Rivas-Torres^{50, 118}, Anand Roopsind⁹², Agustín Rudas⁸⁶, Beatriz Eugenia Salgado Negret⁹³, Masha T. van der Sande³², Flávia Delgado Santana⁹⁴, Flavio Antonio Maës Santos⁵⁷, Rodrigo Scarton Bergamin⁹⁵, Miles R. Silman⁶⁰, Camila Silva⁹⁶, Javier Silva Espejo⁹⁷, Marcos Silveira³⁶, Fernanda Cristina Souza⁹⁸, Martin J. P. Sullivan⁹⁹, Varun Swamy¹⁰⁰, Joey Talbot¹⁰¹, John J. Terborgh¹⁰², Peter van de Meer¹⁰³, Geertje van der Heijden¹⁰⁴, Bert van Uft¹⁰⁵, Rodolfo Vasquez Martinez¹⁰⁶, Laura Vedovato¹⁹, Jason Vleminckx¹⁰⁷, Vincent Antoine Vos⁸², Verginia Wortel¹⁰⁸, Pieter Zuidema³², Joeri Zwerts¹⁰⁹, Susan G. W. Laurance¹¹⁰, William F. Laurance¹¹⁰, Jérôme Chave¹¹¹, James W. Dalling^{112, 113}, Jos Barlow²⁸, Lourens Poorter³², Brian J. Enquist^{114, 115}, Hans ter Steege^{116, 117}, Oliver L. Phillips²⁴, David Galbraith²⁴, Yadvinder Malhi^{1, 120}

Affiliations:

¹Environmental Change Institute, School of Geography and the Environment, University of Oxford; Oxford, OX13QY, UK.

²Consejo Nacional de Investigaciones Científicas y Técnicas, Instituto Multidisciplinario de Biología Vegetal (IMBIV); Córdoba, X5016GCN, Argentina.

³Facultad de Ciencias Exactas, Físicas y Naturales, Universidad Nacional de Córdoba; Córdoba, 5000, Argentina.

⁴School of Biological Sciences, University of Adelaide; Adelaide, 5005, Australia

51 ⁵Facultad de Ciencias Forestales y Ambientales, Universidad Juárez del Estado de
52 Durango, Durango, Mexico
53 ⁶Escuela Politécnica Superior de Ingeniería. Campus Terra. Universidad de Santiago de
54 Compostela, 27002 Lugo, España.
55 ⁷Colegio de Ciencias y Humanidades. Universidad Juárez del Estado de Durango, 34270 Durango,
56 Mexico.
57 ⁸Programa Ciencias Básicas de la Biodiversidad, Instituto de Investigación de Recursos
58 Biológicos Alexander von Humboldt; Bogotá, Colombia.
59 ⁹Institute for Nature Earth and Energy, Pontificia Universidad Católica del Perú; Lima,
60 15088, Peru.
61 ¹⁰Wildlife Conservation Society; New York, 10460, USA.
62 ¹¹Instituto de Biodiversidade e Florestas da Universidade Federal do Oeste do Pará
63 (UFOPA), Rua Vera Paz, s/n (Unidade Tapajós), Bairro Salé, CEP 68040-255, Santarém,
64 Pará, Brasil.
65 ¹² Programa de Pós Graduação em Ecologia e Conservação, Universidade do Estado de
66 Mato Grosso, Nova Xavantina, Brazil
67 ¹³Universidad UNAD-Colombia, Escuela ECAPMA, Bogotá. Cl. 14 Sur # 14-23
68 ¹⁴Center for Tropical Research, Institute of the Environment and Sustainability, University of
69 California, Los Angeles, CA 90095, United States
70 ¹⁵Projeto Dinâmica Biológica de Fragmentos Florestais, Instituto Nacional de Pesquisas da
71 Amazônia - INPA, Av. André Araújo, 2936, Petrópolis, Manaus, AM, 69067-375, Brazil
72 ¹⁶Universidade Federal do Pará / Instituto de Geociências/Faculdade de
73 Meteorologia/Belém - Pará - Brasil.
74 ¹⁷Center for Environmental Studies and Research, University of Campinas (UNICAMP),
75 Campinas 13083-867, Brazil
76 ¹⁸Instituto Nacional de Pesquisas Espaciais—INPE, São José dos Campos, SP, Brazil
77 ¹⁹University of Exeter, Exeter, UK
78 ²⁰Wageningen Research, Wageningen University & Research, PO Box 47, 6700 AA
79 Wageningen, The Netherlands
80 ²¹UNELLEZ-Guanare, Programa de Ciencias del Agro y el Mar, Herbario Universitario
81 (PORT), Venezuela
82 ²²Departamento de Biologia, Universidade Federal do Amazonas. Av. Rodrigo Octavio,
83 6200, Manaus, AM, 69080-900
84 ²³Institute of Biology, University of Campinas – UNICAMP, Campinas, SP, Brazil
85 ²⁴School of Geography, University of Leeds, Leeds LS2 9JT, U.K.
86 ²⁵Naturalis Biodiversity Center, Leiden, Netherlands
87 ²⁶International Center for Tropical Botany (ICTB) Department of Biological Sciences, Florida
88 International University, Miami, Florida, USA
89 ²⁷Centro de Energia Nuclear na Agricultura. Universidade de São Paulo. Av. Centenário
90 303 13416-000 Piracicaba SP Brasil
91 ²⁸Lancaster Environment Centre, Lancaster University, Lancaster, UK
92 ²⁹CIRAD, UPR Forêts et Sociétés, F-34398 Montpellier, France
93 ³⁰Forêts et Sociétés, Univ Montpellier, CIRAD, Montpellier, France
94 ³¹Université de Lorraine, AgroParisTech, INRAE, UMR Silva, 54000 Nancy, France
95 ³²Forest Ecology and Forest Management Group, Wageningen University & Research,
96 Wageningen, The Netherlands
97 ³³Plant Ecology Lab, Federal University of Rio Grande do Sul, Porto Alegre, Brazil
98 ³⁴Woodwell Climate Research Center, 149 Woods Hole Rd. Falmouth, MA, USA
99 ³⁵Centro de Pesquisa Agroflorestal de Roraima, Embrapa Roraima, Boa Vista, Brazil
100 ³⁶Centro de Ciências Biológicas e da Natureza, Universidade Federal do Acre, Campus
101 Universitário, BR 364, Km 04, Distrito Industrial, Rio Branco, AC 69920-900, Brazil

102 ³⁷Department of Forestry, University of Brasilia, Campus Darcy Ribeiro, Brasília 70.900-910,
103 Brazil
104 ³⁸National Park Service, Fredericksburg, VA, USA
105 ³⁹Smithsonian Institution, Washington, DC, USA
106 ⁴⁰Proyecto Castaña, Madre de Dios, Peru
107 ⁴¹Instituto Federal de Educação, Ciência e Tecnologia do Acre, Campus Baixada do Sol,
108 Rua Rio Grande do Sul, 2600 - Aeroporto Velho, Acre, 69911-030
109 ⁴²Pontifícia Universidade Católica do Paraná, Campus Toledo. Rua da União, 500 - Vila
110 Becker. 85902-532 - Toledo PR - Brasil
111 ⁴³Biodiversity and Ecosystem Services, Instituto Tecnológico Vale, Belém, PA, Brazil
112 ⁴⁴Programa de Pós-Graduação em Botânica, Instituto Nacional de Pesquisas da Amazônia
113 (INPA), Manaus, Amazonas 69060-001, Brazil;
114 ⁴⁵Luz da Floresta, Boa Vista, Roraima 69306-320, Brazil.
115 ⁴⁶Instituto de Investigaciones de la Amazonia Peruana, Iquitos, Peru
116 ⁴⁷Universidad Nacional de la Amazonia Peruana, Iquitos, Peru
117 ⁴⁸Cirad, UMR EcoFoG (AgroParistech, CNRS, INRAE, Université des Antilles, Université de
118 la Guyane), Campus Agronomique, Kourou, French Guiana
119 ⁴⁹Department of Anthropology, The University of Texas at Austin, 2201 Speedway Stop
120 C3200, Austin, TX 78712, USA
121 ⁵⁰Estación de Biodiversidad Tiputini, Colegio de Ciencias Biológicas y Ambientales, Universidad
122 San Francisco de Quito (USFQ), Quito, Ecuador.
123 ⁵¹University College London, Dept. of Geography, Gower Street, London, WC1E 6BT, U
124 ⁵²NERC National Centre for Earth Observation (NCEO), Gower Street, London, WC1E 6BT,
125 UK.
126 ⁵³Biology Department and Sabin Center for Environment and Sustainability, Wake Forest
127 University, Winston-Salem, NC, USA
128 ⁵⁴Herbario Vargas (CUZ), Escuela Profesional de Biología, Universidad Nacional de San
129 Antonio Abad del Cusco, Cusco, Peru
130 ⁵⁵School of Geography, Earth and Environmental Sciences, University of Plymouth,
131 Plymouth, UK
132 ⁵⁶Geography, Faculty of Environment, Science, and Economy, University of Exeter, Exeter,
133 UK
134 ⁵⁷Department of Plant Biology, Institute of Biology, University of Campinas – UNICAMP, CP
135 6109, Campinas, SP, 13083-970, Brazil
136 ⁵⁸Department of Natural Sciences, Maths, and Education, Centre for Agrarian Sciences,
137 Federal University of São Carlos (UFSCar), Rodovia Anhanguera - SP 330, km 174, Araras,
138 SP 13600-970, Brazil
139 ⁵⁹AMAP (Botanique et Modélisation de l'Architecture des Plantes et des Végétations),
140 Université de Montpellier, CIRAD, CNRS, INRAE, IRD, Montpellier, France
141 ⁶⁰Biology Department and Sabin Center for Environment and Sustainability, Wake Forest
142 Univ., Winston-Salem, NC, USA
143 ⁶¹Federal University of Acre, Cruzeiro do Sul, Brazil
144 ⁶²Forêts et sociétés, Univ Montpellier, CIRAD, Montpellier, France
145 ⁶³Instituto Venezolano de Investigaciones Científicas (IVIC), Caracas, Venezuela
146 ⁶⁴Royal Botanic Gardens, Kew, London, Richmond, TW9 3AE, UK
147 ⁶⁵Herbario Alwyn Gentry (HAG), Universidad Nacional Amazónica de Madre de Dios
148 (UNAMAD), Av. Jorge Chávez 1160. Puerto Maldonado, Madre de Dios, Perú.
149 ⁶⁶Centro Ecológico INKAMAZONIA, Valle de Kosñipata, vía Cusco-Reserva de Biósfera del
150 Manú. Cusco, Perú.
151 ⁶⁷Broward County Parks & Recreation Division, Oakland Park, FL, 33309 USA

152 ⁶⁸Dept Biological Sciences, Florida Atlantic University, 777 Glades Rd, Boca Raton, FL
153 33431
154 ⁶⁹Grupo de Investigación en Ecología y Conservación de Fauna y Flora Silvestre, Instituto
155 Amazónico de Investigaciones Imani, Universidad Nacional de Colombia – Sede Amazonia,
156 Km. 2 antigua vía Leticia-Tarapacá, Amazonas, Colombia, Suramérica
157 ⁷⁰Unique land use GmbH, Schnewlinstraße 10, 79098 Freiburg im Breisgau, Germany.
158 ⁷¹Smurfit Kappa Colombia, CALLE 15 18-109 Barrio La Estancia, Yumbo, Valle del Cauca,
159 Colombia
160 ⁷²Universidad Nacional de Colombia Medellín, Cra. 65 #59a-110, Medellín, Robledo,
161 Medellín, Antioquia, Colombia
162 ⁷³Instituto Nacional de Pesquisas da Amazônia, Manaus AM Brazil
163 ⁷⁴Laboratory of Phytogeography and Evolutionary Ecology, Department of Forest Sciences,
164 Federal University of Lavras, Lavras, Minas Gerais, Brazil
165 ⁷⁵Universidad del Tolima, Ibagué, Colombia
166 ⁷⁶Plant Ecology and Evolution, Evolutionary Biology Center, Uppsala University, Uppsala,
167 Sweden
168 ⁷⁷Coordenação de Biodiversidade, Instituto Nacional de Pesquisas da Amazônia, Av. André
169 Araújo 2936, Manaus, Amazonas, Brasil
170 ⁷⁸Universidad Estatal Amazónica, Puyo, Pastaza, Ecuador
171 ⁷⁹Herbario Nacional del Ecuador, Universidad Técnica del Norte, Ecuador
172 ⁸⁰Universidad Nacional Agraria La Molina, Av. La Molina s/n, La Molina, Lima.
173 ⁸¹Universidad Nacional de San Antonio Abad del Cusco, Perú
174 ⁸²Instituto de Investigaciones Forestales de la Amazonía, Universidad Autónoma del Beni
175 José Ballivián, Riberalta, Beni, Bolivia.
176 ⁸³College of Life and Environmental Sciences, University of Exeter, Exeter, Devon EX4 4QE,
177 UK
178 ⁸⁴Tropical Diversity Section, Royal Botanic Gardens Edinburgh, Edinburgh EH3 5LR, UK
179 ⁸⁵Science & Education, Field Museum of Natural History, 1400 S. DuSable Lake Shore
180 Drive, Chicago, Illinois 60605 USA
181 ⁸⁶Instituto de Ciencias Naturales, Universidad Nacional de Colombia. Ciudad Universitaria,
182 Carrera 30 No. 45-03 Edif 425. Bogotá. Colombia. CP 111321
183 ⁸⁷Coordination of Environmental Dynamics, National Institute for Amazonian Research,
184 Manaus, Brazil
185 ⁸⁸Instituto de Investigaciones para el Desarrollo Forestal (INDEFOR), Universidad de los
186 Andes, Mérida, Venezuela
187 ⁸⁹AMAP, Univ. Montpellier, IRD, CNRS, CIRAD, INRAE, Montpellier, France
188 ⁹⁰Corporación COL-TREE, Medellín, Colombia
189 ⁹¹Facultad de Ciencias Forestales, Universidad Nacional Agraria La Molina, Lima, Perú, Av.
190 La Molina s.n., La Molina, Lima
191 ⁹²Conservation International, Arlington, United States
192 ⁹³Departamento de Biología, Universidad Nacional de Colombia, Bogotá, Colombia. Carrera
193 45 N° 26-85
194 ⁹⁴Instituto Nacional de Pesquisas da Amazônia (INPA), Av. André Araújo, 2936 - Petrópolis,
195 Manaus - AM, 69067-375
196 ⁹⁵Birmingham Institute of Forest Research (BIFoR), University of Birmingham
197 ⁹⁶Instituto de Pesquisa Ambiental da Amazônia (IPAM), SCLN 211, Bloco B, Sala 201,
198 Bairro Asa Norte, Brasília-DF | 70863-520
199 ⁹⁷Departamento de Biología. Universidad de La Serena. Chile
200 ⁹⁸Departamento de Ecologia e Conservação, Instituto de Ciências Naturais, Universidade
201 Federal de Lavras, Lavras, Minas Gerais, Brazil.
202 ⁹⁹Department of Natural Sciences, Manchester Metropolitan University, Manchester, UK

203 ¹⁰⁰Center for Energy, Environment & Sustainability, Wake Forest University, USA
204 ¹⁰¹Institute for Transport Studies, University of Leeds, Leeds, UK
205 ¹⁰²Department of Biology, University of Florida, Gainesville, Florida 32611
206 School of Science and Engineering, James Cook University, Cairns, Queensland, Australia
207 ¹⁰³Van Hall Larenstein University of Applied Sciences, Larensteinselaan 26a P.O.Box 9001,
208 6880 GB Velp, The Netherlands
209 ¹⁰⁴School of Geography, University of Nottingham, University Park, Nottingham, NG7 2RD,
210 UK
211 ¹⁰⁵Royal Netherlands Meteorological Institute (KNMI), PO box 201, 3730AE De Bilt, the
212 Netherlands
213 ¹⁰⁶Jardín Botánico de Missouri, Prolongación Bolognesi Mz. E-6 Oxapampa-Perú
214 ¹⁰⁷Université Libre de Bruxelles. 50 Av F.D. Roosevelt, 1050 Brussels
215 ¹⁰⁸Department of Forest Management, Centre for Agricultural Research in Suriname,
216 CELOS. Prof.Dr.Ir.J.Ruinardlaan #1, Paramaribo. Suriname
217 ¹⁰⁹Utrecht University, Padualaan 8, 3584 CH Utrecht, The Netherlands.
218 ¹¹⁰Centre for Tropical Environmental and Sustainability Science, College of Science and
219 Engineering, James Cook University, Cairns, Queensland, Australia
220 ¹¹¹Laboratoire Evolution et Diversité Biologique, CNRS, Université Paul Sabatier, IRD,
221 UMR; Toulouse, 5174 EDB, France.
222 ¹¹²Department of Plant Biology, University of Illinois, Urbana-Champaign; IL, 61801, USA
223 ¹¹³Smithsonian Tropical Research Institute; Ancon, Republic of Panama
224 ¹¹⁴Department of Ecology and Evolutionary Biology, University of Arizona; Tucson, AZ
225 85721, USA.
226 ¹¹⁵Santa Fe Institute; 1399 Hyde Park Rd., Santa Fe, NM 87501, USA.
227 ¹¹⁶Tropical Botany, Naturalis Biodiversity Center, Leiden, The Netherlands
228 ¹¹⁷Quantitative Biodiversity dynamics, Department of Biology, Utrecht University, Utrecht,
229 The Netherlands
230 ¹¹⁸Department of Wildlife Ecology and Conservation, University of Florida, 110285 Newins-
231 Ziegler Hall, Gainesville, FL, 32611, USA.
232 ¹¹⁹Departamento de Ciencias Forestales, Facultad de Ingeniería Forestal, Universidad del
233 Tolima. Colombia
234 ¹²⁰Leverhulme Centre for Nature Recovery, University of Oxford
235
236
237
238
239 * Corresponding author Email: jeaggu@gmail.com

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241 **Abstract:** Understanding ecosystem responses to climate change is among the greatest challenges
242 in biosphere science. This challenge is particularly acute in high biodiversity tropical forests. Here
243 we examine how the functional composition of tropical forests of the Americas has shifted in recent
244 decades. We found that functional shifts in the tree assemblages are insufficient to keep track with
245 what would be expected given observed climate change, typically changing less than 8% of the
246 required amount because of intrinsic inertia in tree demography. However, recruiting assemblages
247 show stronger shifts of 21% of the proportional climate change expectation. Overall, trees of the
248 most diverse forests on Earth are changing in functional trait composition, but at a rate that is
249 insufficient to track climate change.

250

251 **One-Sentence Summary:** The composition of tropical forests in the Americas is changing but not
252 fast enough to keep track of climate change.

253 **Main Text**

254 Forest responses to human-driven perturbations like climate change will determine the diversity and
255 function of the terrestrial biosphere through this century and beyond. Here we focus on tropical
256 American forests (henceforth referring to all forests encompassing continental areas from Brazil to
257 Mexico). These host the greatest concentration of tree species in the world (1), including six key
258 biodiversity hotspots (2) and half of Earth's most intact tropical forests (3). In the face of threats from
259 climate change and continuing loss in area and integrity (3, 4, 5), it is both critical and urgent to
260 understand the ability of these complex systems to adapt to change and survive.

261 The relationships between environmental conditions, plant performance, and distribution is
262 mediated by species' functional traits (6, 7, 8). These morphological, structural, chemical, and
263 phenological characteristics show consistent relationships with climate and other drivers, such as
264 soil conditions and biotic interactions (9). Notably, recent work has shown positive relationships
265 between mean annual temperature and leaf area, specific leaf area, leaf nitrogen, wood density,
266 seed mass (10), and leaf thickness (11). Other work has detected a negative relationship with
267 elevation for specific leaf area, and leaf nitrogen (11). Hence, a traits-based approach provides a
268 promising framework for predicting the impacts of climate change and resilience across forest
269 ecosystems (7, 12, 13). Such a framework can be used to understand responses of key ecosystem
270 functions, such as forest carbon storage responses to climate change and extreme events (14, 15).

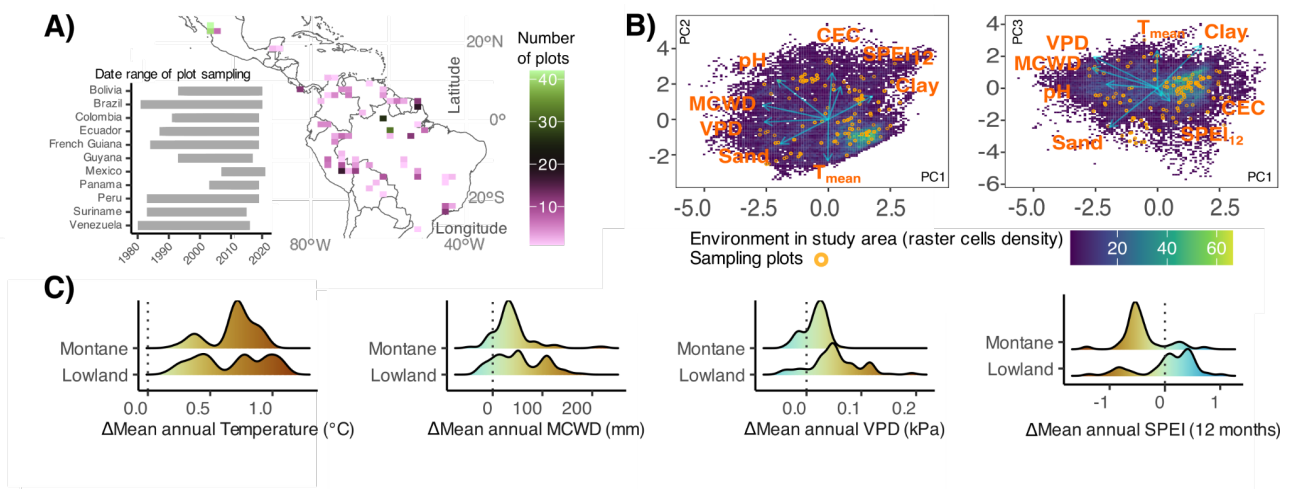
271 Tropical land regions are experiencing rapid climate change (16), which is already affecting
272 tropical American plant communities. In Amazonia, changes in precipitation patterns and more
273 frequent droughts have led to an increase in the recruitment of dry-affiliated species (xerophilization)
274 (17). In the Andes, rising temperatures have led to increasing abundances of species tolerant to
275 higher temperature (thermophilization) (18). Across Mesoamerica it is expected that climate change
276 will cause an expansion of tropical dry forests to higher elevations (over 200 m above current
277 average elevation) (19). However, tree species may be unable to shift their distribution fast enough
278 to track their climatic niche, given their slow demography and the prevalence of dispersal limitation
279 and competitive abilities (20), with such responses also being determined by environmental
280 tolerances at different life stages (21). Moreover, recent work has shown large range contractions of
281 tree species rather than range shifts (22). Changes in climate across the tropical Americas are
282 expected to become stronger, with some scenarios projecting temperature increases of up to ~4°C
283 and precipitation reductions of close to 20% by 2100 (23, 24, 25). This would expose current species
284 assemblages to climates they have not previously experienced (26), potentially selecting for future
285 plant communities adapted to such climates but unlike those currently observed (27). Community
286 responses to climate change will thus likely depend on underlying mechanisms and geographical
287 context.

288 If species respond to climate change via migration, we would expect montane communities
289 to track climate change better than those in the lowland forests because different climate conditions
290 occur at shorter distances and hence are potentially easier to migrate to than in the lowlands (28,
291 29, 30, 31). Given exposure to a drying and warming climate, we could reasonably expect increased
292 abundance of species exhibiting more drought-tolerance traits (i.e. in the 'slow' plant economics
293 spectrum) (32), such as high wood density (e.g. to prevent cavitation) (33) and smaller, thicker leaves
294 (e.g. for lower evapotranspiration and reduced radiation exposure) (34). Other drought-avoidance
295 traits, notably deciduousness (often associated with more acquisitive leaves), could also become
296 more prominent in the future as an adaptation to increasing drought conditions and to past
297 disturbances (35, 36). Seed traits play a pivotal role in the reproduction and dispersal, or migration
298 capacity of species (37). Under a warming and drying climate we might expect species with smaller
299 seeds to increase in abundance (38). This is because wind-dispersed seeds, most frequent in drier
300 and more seasonal biomes, tend to be smaller than animal-dispersed seeds (39). However, other
301 factors, such as defaunation of frugivorous seed-dispersing large mammals and birds, may more
302 strongly drive their shifts at short time scales (40). We would also expect montane forests to show
303 stronger functional responses to climate change than lowland forests given their more varied climatic

304 conditions at shorter distances (18, 41) and due to differences in soil nutrient availability (e.g., N:P
 305 ratios) across elevation gradients (42).

306 It is still unclear how shifts in the abundance and distribution of species translate into changes
 307 in functional composition, and what functional changes have occurred through the last half century
 308 as a response to the onset of a warmer, drier and more variable climate across the tropical Americas.
 309 Moreover, it is unknown if forest level functional shifts are more attributable to the surviving tree
 310 individuals or to the addition (i.e. recruitment) or removal (i.e. mortality) of trees to the assemblage.
 311 It is also uncertain if these functional shifts match the direction of climate change, and if so, whether
 312 the rate of functional trait change keeps pace with climate change or lags behind.

313 Here, we address these knowledge gaps by monitoring and analysing 415 long-term forest
 314 plot sites spanning more than 40 years (1980 - 2021) and encompassing >250,000 individual trees
 315 across the tropics from Mexico to southern Brazil. In these sites, tree diversity, structure and function
 316 were observed on the ground over the long term. Our effort spans structurally intact forests from the
 317 lowland tropics (hereafter forest plots <700 m elevation, see methods) to pre-montane and montane
 318 forests (>700 m elevation; henceforth referred to as montane) from the Andes to subtropical fringes
 319 (Fig. 1; Table S1). We combine this monitoring and analysis of changes in the plant community
 320 composition dataset with measurements and detailed assessment of 12 plant functional traits
 321 involved in responses to a changing climate. These include photosynthetic capacity (A_{sat}), leaf
 322 chemistry (content of carbon: C, nitrogen: N and phosphorus: P), leaf area (Area), specific leaf area
 323 (SLA), leaf fresh mass (FM), leaf thickness (Thickness), abundance of deciduous species (DE),
 324 maximum species tree height (H_{max}), wood density (WD) and seed mass (SM) (Table S2). We first
 325 investigate current plant trait-environment relationships for lowland and montane forests. We
 326 examine how and where these ecosystems have shifted in their functional trait composition due to
 327 changes in the plant community taxonomic composition. Because of the long lifespan of tropical
 328 trees (43) and their slow turnover, we performed this analysis at the whole community level, for
 329 survivors, and for the recruiting and mortality assemblages, with the last two expected to show faster
 330 change. We also analysed if observed changes in trait composition have been enough to track
 331 climate change to date. This climate change tracking analysis was carried out for the full community,
 332 survivors and the recruitment assemblages and not for mortality because these individuals will not
 333 contribute to future change. We hypothesise that for most traits there will be a consistent trait-climate
 334 relationship but this will vary according to functional trait and forest type, and that lowland and
 335 montane forests will differ in their functional responses to climate change given their different
 336 exposures to climatic conditions with spatial distance. We expect that, given the long lifespan of
 337 tropical trees and rapid pace of recent climate change, forests will demonstrate ecological inertia, so
 338 that changes in functional composition lag behind changes in climate.



339 **Fig. 1. Study area showing the distribution and number of vegetation plots sampled across time (A), principal**
 340 **component analysis (PC1, PC2 and PC3) depicting the climate and soil chemistry and texture space available in**
 341 **the study area (T_{mean} : mean air temperature, MCWD: maximum climatic water deficit, $SPEI_{12}$: standardised**
 342

343 precipitation-evapotranspiration index, VPD: vapour pressure deficit, CEC: soil cation exchange capacity, soil pH,
344 sand and clay amount) and the location of the sampling plots in the environmental space (B), and change in
345 climate conditions (1980-1990 vs 2010-2020) in the plot network (C). In B) PC1 is mainly loaded by the maximum
346 climatic water deficit (MCWD: -0.527) and Vapour Pressure Deficit (VPD: -0.515), PC2 by air temperature (T_{mean} : -0.465)
347 and soil cation exchange capacity (CEC: 0.524) and PC3 by soil clay % (-0.535) and soil sand % (0.486). In C) the vertical
348 dotted lines indicate zero change. Brown colours depict increases in temperature, drier conditions (for MCWD and VPD)
349 or increased drought intensity (for SPEI: standardised precipitation evapotranspiration index). Blue colours depict an
350 increase in water availability. In MCWD larger positive values indicate higher water stress. Climate data was derived from
351 the TerraClimate project (44) and soil data from SoilGrids.org.
352

353 **Current trait-environment relationships**

354 To evaluate contemporary trait-climate relationships across tropical American forests, we used data
355 from 415 forest plots (mean plot size 0.88 [min: 0.12, max: 25] ha and 5.7 [min: 2, max: 41] censuses
356 per plot), for which we extracted climate (44) and soil (SoilGrids.org) data for their sampling years
357 (1980-2021). As the most dominant species are expected to drive ecosystem processes (45), for
358 each plot we calculated the community-weighted mean of each plant functional trait (Table S2;
359 Methods) based on the relative basal area of the species and their trait value (hereafter “community
360 functional traits”). We then modelled each community functional trait as a function of the additive
361 effects of relevant and largely uncorrelated climatic drivers of species distributions (Fig. S1), i.e., the
362 mean annual values (between 1980-2021) of temperature (T_{mean}), vapour pressure deficit (VPD_{mean})
363 (46), maximum climatic water deficit ($\text{MCWD}_{\text{mean}}$, here with positive values depicting higher water
364 stress) (47) and standardised precipitation-evapotranspiration index (SPEI_{12}) (48), each one of these
365 interacting with forest type (lowland or montane) (Methods). In order to control for possible soil
366 effects which impact plant distributions (13), we included the additive effect of cation exchange
367 capacity (CEC), pH, and the percentage of clay and sand for each plot location in the models
368 (Methods).
369

370 Several community functional traits show consistent relationships with climate, with
371 temperature showing some of the strongest effects (Table S3; Fig. S2). For both lowland and
372 montane forest types, an increase in temperature (T_{mean}) across space is associated with an increase
373 in community-mean leaf area and seed mass, and a decrease in photosynthetic capacity, specific
374 leaf area, and the proportion of deciduous species. However, the geographical relationship across
375 lowland and montane forests is not consistent for leaf chemistry (leaf carbon, nitrogen and
376 phosphorus content), wood density, maximum species tree height, leaf fresh mass or leaf thickness
377 (Fig. S2). An increase in water stress ($\text{MCWD}_{\text{mean}}$) is associated with an increase in photosynthetic
378 capacity, leaf nitrogen content, leaf area and wood density across lowland forests and decreases in
379 montane forests, with decreases in specific leaf area and maximum species tree height for both
380 forest types (Table S3; Fig. S2). The increase in these leaf traits in drier forests could be associated
381 with the high photosynthetic rates generally attained by deciduous species over the growing season
382 (49, 50) and the fact that lower maximum tree height and higher wood density tend to correlate with
383 higher resistance to lethally low levels of soil moisture availability (51). However, consistent climatic
384 relationships across both forest types are not apparent for the other traits analysed (Table S3; Fig.
385 S2). One plausible explanation is that this reflects their different position along the climatic gradient
386 (i.e. temperature and precipitation), with lowlands occupying areas with more homogeneous climate
387 conditions across large spatial extents in comparison to montane forests, which span a large range
388 of climates across smaller spatial extents. Alternatively, the differences between lowland and
389 montane forests are potentially due to additional variables, such as cloud immersion effects in upper
390 montane forests, which could shift the nature of trait-environment relationships given the lower
391 radiation and temperature, and higher water availability across the year (52, 53, 54).

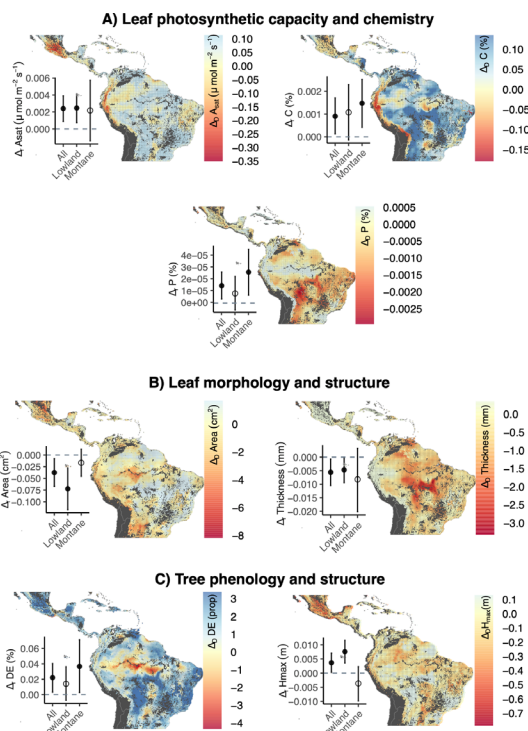
392 **Changes in trait composition across time**

393 We next analysed if and how the functional trait composition of tropical American forests has shifted,
394 and how much of this can be explained by observed changes in climate over the past 40 years
395 (Methods). We first calculated the community-weighted mean (CWM) of each plant functional trait
396 for each vegetation census available for the full live tree community (i.e. including all trees alive in a
397 given census), henceforth ‘full community assemblage’, and separately for the survivor tree
398 assemblage (individuals that are alive in two subsequent censuses, e.g. from census one to census
399 two), tree recruitment (individuals not present in the previous census and recruited in the subsequent

400 census) and tree mortality (individuals alive in previous census but dead in the subsequent census)
401 assemblages. We define the recruiting assemblage as individuals that passed the threshold of 10
402 cm DBH in a given census and were not present in the previous census. Then we calculated their
403 yearly rate of change across time (Methods). We tested if the changes in trait CWM differed from
404 zero when using all vegetation plots together and when separated into lowland and montane forests.
405 We calculated the Highest Density Interval (HDI) containing the 95% most probable effect values
406 and considered it significant when the HDI did not overlap 0. We then investigated whether the
407 observed shifts in trait CWM differed significantly between lowland and montane forests (Methods).
408 For shorthand and readability, all mention of mean traits and shifts below refer to CWM trait values.

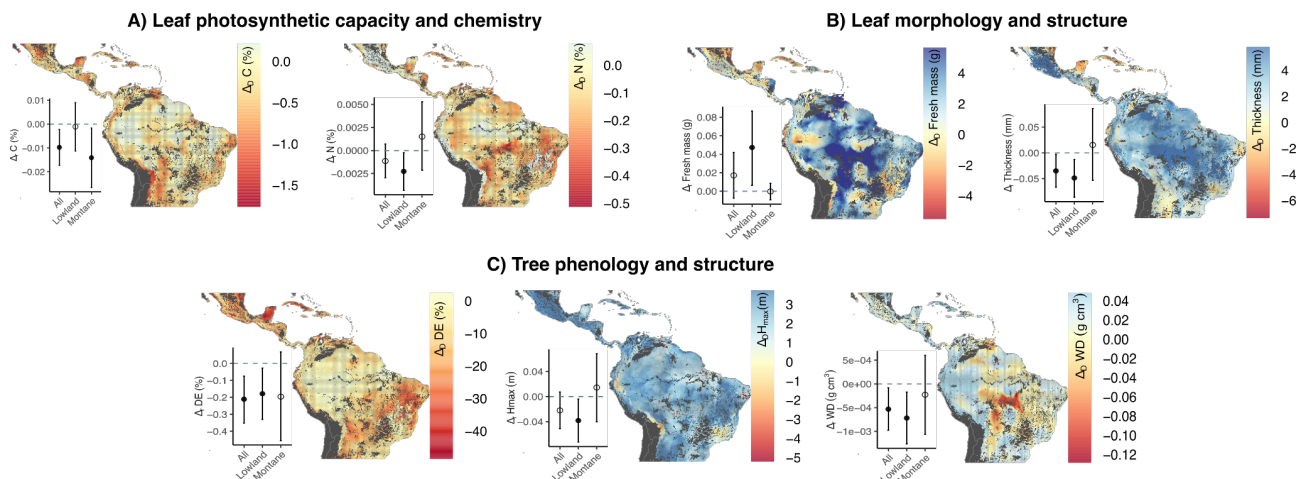
409 When considering all plots together for the full community assemblage, we found that seven
410 out of the 12 traits analysed exhibited significant changes in their CWM values (insets in Fig. 2). Only
411 leaf nitrogen, fresh mass, specific leaf area, seed mass and wood density did not show significant
412 shifts across time (Table S4; Fig. S3). The survivors assemblage showed the same pattern of
413 community trait changes (Table S4; Fig. S4) as the full community assemblage with the main
414 differences being a significant decrease in leaf fresh mass in the lowlands for the survivors
415 assemblage. Hence, hereafter we focus on the results from the full community, recruiting and
416 mortality assemblages. From the community traits with significant changes, we found an average
417 increase in photosynthetic capacity of $0.0024 \mu\text{mol m}^{-2} \text{s}^{-1} \text{year}^{-1}$ (HDI-low and HDI-high: 0.0009,
418 0.0039), leaf carbon content $0.0009\% \text{year}^{-1}$ (0.0001, 0.0017), phosphorus $1.5 \times 10^{-5}\% \text{year}^{-1}$
419 (3.58×10^{-6} , 2.63×10^{-5}), the abundance of deciduous species $0.02\% \text{year}^{-1}$ (0.003, 0.04) and species
420 maximum height $0.0037 \text{m year}^{-1}$ (0.0002, 0.007), while community leaf area decreased on average
421 $-0.03 \text{cm}^2 \text{year}^{-1}$ (-0.06, -0.007) and leaf thickness decreased $-0.005 \text{mm year}^{-1}$ (-0.01, -0.0007)
422 (insets in Fig. 2; Table S4). In the lowland forests, we detected significant trait changes for four
423 (increasing: photosynthetic capacity and species maximum height; decreasing: leaf area and
424 thickness) out of the 12 traits analysed (Table S4; insets in Fig. 2). Montane forests showed
425 significant increases in leaf carbon, phosphorus and the abundance of deciduous species (Table
426 S4; Fig. 2 insets).

427 **Fig. 2. Estimated changes in mean community functional trait values across time for tropical American forests.**
428 Only traits with significant changes are shown, with others shown in Fig. S3. A) Changes in trait community-weighted mean
429 (CWM) for leaf photosynthetic capacity and leaf chemistry traits, B) for leaf morphology and structural traits and C) for tree
430 phenology and structural traits. The insets in the left-hand side of each map show the observed yearly rate of change,
431 obtained from sampled vegetation plots, from the statistical models in Table S4 for all forests together and only for lowland
432 or montane forests. In the A-C insets, significant shifts are shown as filled circles and non-significant as empty circles. The
433 vertical lines depict the Highest Density Intervals (95% HDI), and the horizontal grey dotted line indicates zero change.
434 Maps show the decadal predicted changes (ΔD) in trait CWM across tropical American forests given changes in climate
435 derived from the statistical model fits with R-squared values ranging between 0.21 and 0.34 (Table S5). In the maps,
436 warmer colours represent decreases in the trait CWM, and cooler colours increase in the trait CWM, with yellow-white
437 colours representing slight or no change. The grey mask on the background of each map represents all predominately
438 non-forested areas (e.g., crop fields, swamps, open woodland, areas with small patches of trees, deserts and alpine
439 regions) and was derived from the European Space Agency Land Cover CCI Product (81). A_{sat} : photosynthetic capacity at
440 light saturation, C: leaf carbon content, P: leaf phosphorus content, Area: leaf area, Thickness: leaf thickness, DE:



441 deciduousness, H_{max} : adult maximum height. R-squared values of community functional trait model predictions range from
 442 21%-34% and are shown in Table S5.

443 The recruitment assemblage experienced significant changes for seven traits, with six
 444 showing decreases, i.e. leaf carbon content $-0.014\% \text{ year}^{-1}$ ($-0.02, -0.001$; in montane forests), leaf
 445 nitrogen content $-0.002\% \text{ year}^{-1}$ ($-0.004, -0.0002$), leaf thickness $-0.04 \text{ mm year}^{-1}$ ($-0.08, -0.01$),
 446 deciduousness $-0.17 \text{ % year}^{-1}$ ($-0.33, -0.02$), tree adult maximum height ($-0.03 \text{ m year}^{-1}$ [$-0.07, -$
 447 0.003], and WD: $-0.0007 \text{ g cm}^3 \text{ year}^{-1}$). The leaf fresh mass of recruits increased on average 0.04 g
 448 year^{-1} ($0.006, 0.08$; all but leaf carbon in lowland forests; insets in Fig. 3; Table S4).

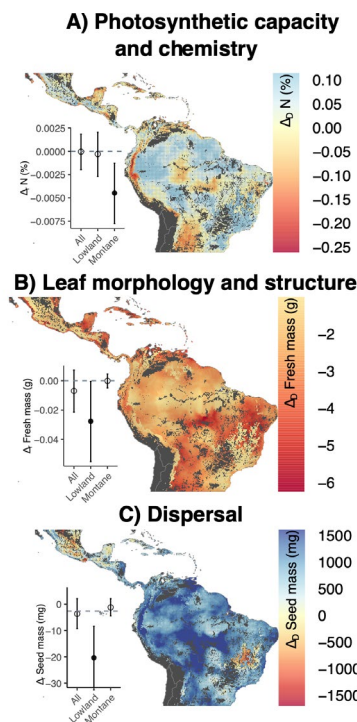


449
 450 **Fig. 3. Estimated changes in mean community functional trait values across time for tropical American forests for**
 451 **the recruitment assemblage.** Only traits with significant changes are shown, with others shown in Fig. S5. A) Changes
 452 in trait community-weighted mean (CWM) for the recruitment assemblage for leaf chemistry traits, B) for leaf morphology
 453 and structural traits, and C) for tree phenology and structural traits. There were no significant changes recorded for
 454 dispersal traits. The insets in the left-hand side of each map show the observed yearly rate of change, obtained from
 455 sampled vegetation plots, from the statistical models in Table S4 for all forests together and only for lowland or montane
 456 forests. In the A-C insets, significant shifts are shown as filled circles and non-significant as empty circles. The vertical
 457 lines depict the Highest Density Intervals (95% HDI), and the horizontal grey dotted line indicates zero change. Maps show
 458 the decadal predicted changes (ΔD) in trait CWM across tropical American forests given changes in climate as a result
 459 of the statistical model fits with R-squared values ranging between 0.43 and 0.65 (Table S5). In the maps, warmer colours
 460 represent decreases in the trait CWM, and cooler colours increase in the trait CWM, with yellow-white colours representing
 461 slight or no change. The grey mask on the background of each map represents all predominately non-forested areas (e.g.,

462 crop fields, swamps, open woodland, areas with small patches of trees, deserts and alpine regions) and was derived from
 463 the European Space Agency Land Cover CCI Product (81). C: leaf carbon content, N: leaf nitrogen content, Fresh mass:
 464 leaf fresh mass, Thickness: leaf thickness, DE: deciduousness, H_{max}: adult maximum height, WD: wood density.

465 For the mortality assemblage, only the CWM of leaf nitrogen content $-0.004\% \text{ year}^{-1}$ (-0.007 ,
 466 -0.001 ; montane forests), leaf fresh mass, $-0.02 \text{ g year}^{-1}$ (-0.05 , -0.0003) and seed mass -17.7 mg
 467 year^{-1} (-29.9 , -5.7) in lowland forests experienced significant declines (insets in Fig. 4; Table S4).

468 To help identify the underlying climatic drivers of forest functional change, we used
 469 multivariate linear models to estimate the yearly change (Δr ; i.e. from first to last census), in the trait
 470 values (Δ_r trait CWM) as a function of the yearly rate of change in temperature (ΔT_r), maximum
 471 climatic water deficit (ΔMCWD_r), standardised precipitation-evapotranspiration index (ΔSPEI_r) and
 472 vapour pressure deficit (ΔVPD_r), each one of these interacting with forest type, and accounted for
 473 soil characteristics by including in the models the CEC, pH, clay and sand content (Methods; maps
 474 in Fig. 2, Fig. 3 and Fig. 4). Our results for the full community assemblage, survivors and for
 475 recruitment and mortality assemblages (Table S5) demonstrate the role of climate, specifically
 476 temperature and water availability, as a determinant of trait shifts across the forests, and show the
 477 differences in response between lowland and montane forests (Table S5). Our mapped model
 478 predictions (Fig. 2, Fig. 3 and Fig. 4 maps) depict in a spatially explicit way areas where stable CWM
 479 trait values (light yellow and light blue), their increases (darker blue) or decreases (yellow to red) are
 480 predicted to have occurred across tropical American forests with some of the strongest CWM trait
 481 shifts predicted across forests in Amazonia.



482 **Fig. 4. Estimated changes in mean community functional trait values across time for tropical American forests for**
 483 **the mortality assemblage.** Only traits with significant changes are shown, with others shown in Fig. S6. A) Changes in
 484 trait community-weighted mean (CWM) for the mortality assemblage for leaf chemistry traits, B) for leaf morphology and
 485 structural traits, and C) for dispersal traits. There were no significant changes recorded for tree phenology and structural
 486 traits. The insets in the left-hand side of each map show the observed yearly rate of change, obtained from sampled
 487 vegetation plots, from the statistical models in Table S4 for all forests together and only for lowland or montane forests. In
 488 the A-C insets, significant shifts are shown as filled circles and non-significant as empty circles. The vertical lines depict
 489 the Highest Density Intervals (95% HDI), and the horizontal grey dotted line indicates zero change. Maps show the decadal
 490 predicted changes (ΔD) in trait CWM across tropical American forests given changes in climate as a result of the statistical
 491 model fits with R-squared values ranging between 0.43 and 0.65 (Table S5). In the maps, warmer colours represent
 492 decreases in the trait CWM, and cooler colours increase in the trait CWM, with yellow-white colours representing slight or
 493 no change. The grey mask on the background of each map represents all predominately non-forested areas (e.g., crop
 494 fields, swamps, open woodland, areas with small patches of trees, deserts and alpine regions) and was derived from the

495 European Space Agency Land Cover CCI Product (81). N: leaf nitrogen content, Fresh mass: leaf fresh mass, Seed mass:
496 seed mass weight.

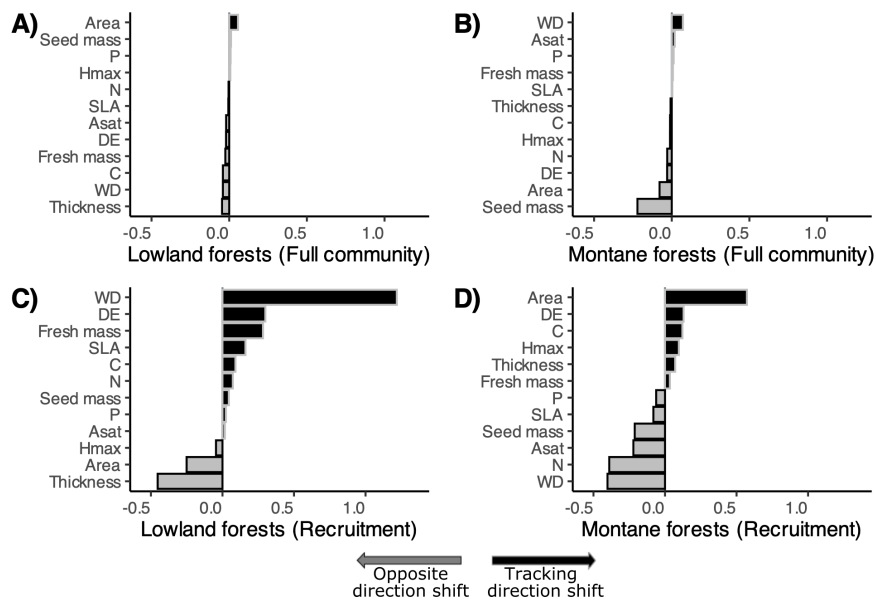
497 **Can tropical American forest functional composition track climate change?**

498 We next examined whether the observed community trait changes are sufficient to maintain
499 expected trait-environment relationships for the full community, the survivors, and the recruiting
500 assemblages. We expected recruitment to be more sensitive to climate change as the full community
501 is dominated by the demographic inertia of established adult trees. To quantify the trait changes that
502 would be necessary for forest communities to track predicted climate change, we first quantified the
503 relationship between community traits and environment before most anthropogenic climate changes
504 occurred (1980-2005; i.e., as baseline CWM trait-environment relationships). We took our observed
505 trait-climate relationships (built with the 1980-2005 period data; Table S6) and used them to predict
506 the trait CWM to the 1980-2005 climate conditions plus the observed changes in climate across the
507 study sites for the full time period (the last 40 years; Methods; Table S4). This allowed us to predict
508 the CWM trait values that the forests should have if they fully tracked recent climate change (Tables
509 S7 and S8). The ratio between the observed and the expected changes (for the full and the recruiting
510 assemblages) indicates how well these forest traits are tracking these climate equilibrium predictions
511 based on community changes alone (Methods; Fig. 5).

512 Our results show that for all measured traits of the full community and survivor assemblages,
513 tree species community composition is not changing sufficiently to track climate change but that the
514 recruiting community shows the largest shifts (Fig 5; results for survivors and mortality assemblages
515 in Fig. S7). At the region-wide scale for the full community assemblage, all traits show less than 6%
516 (average of 1.9% of positive shifts) for lowland forests and 7% (average of 2% of positive shifts) for
517 montane forest of the change required. Several traits show very little change or even modest
518 changes in the opposite direction to those expected (Fig. 5A and Fig. 5B; Table S7 and Table S8).
519 For the survivor assemblages, all traits show less than 8% (average of 2.8% of positive shifts) for
520 lowland forests and 4% (average of 1.7% of positive shifts) for montane forest of the change required
521 (Fig. S7). We detect larger community trait shifts in the recruiting assemblages of an average 21.8%
522 of the change required for lowland forests and 17.5% for montane forests when only positive tracking
523 traits are considered. When both positive and negative tracking is considered, the recruitment
524 assemblage shows an average tracking of 11.4% for lowland and -0.67% for mountain forests (Fig.
525 5C and Fig. 5D; Table S7 and Table S8). In lowland forests, community mean wood density appears
526 to be changing fast enough in the recruiting assemblages to track climate change expectation.
527 Overall, we see some evidence of how the recruiting forest assemblages of lowland and montane
528 forests are shifting their community traits, often for different sets of community mean trait values, in
529 response to climate change. However, for most traits even the recruiting community does not seem
530 to be changing quickly enough to track climate change. More significant community trait shifts have
531 occurred in lowland than in montane forests, which is consistent with a more rapidly drying climate
532 in lowland forests (Fig. 2, Fig. 3 and Fig. 4; Table S4).

533

534 **Fig. 5. Tracking of trait community weighted mean (CWM) for the full live (A, B) and recruiting (C, D) assemblages**
535 **in lowland (A, C) and montane (B, D) forests given the observed changes in climate across the sampling plots.**
536 The X axis shows the ratio of changes in trait CWM, based on actual trait CWM changes observed at the plot level through
537 time, versus expected changes in trait CWM, based on spatial climate-trait relationships given observed changes in
538 climate. Positive values (black bars) indicate that observed and predicted changes are both positive or both negative and,
539 hence, are going into the same direction, whereas negative values (grey bars) indicate that observed and predicted
540 changes are going in opposite directions. A ratio of change value of one would indicate perfect tracking. The Y axis shows
541 the traits sorted by the change ratio amount (see full statistical details in Table S7 and Table S8). Values of zero and close
542 to zero represent no or slight trait shifts. A_{sat} : photosynthetic capacity at light saturation, C: leaf carbon content, N: leaf



543 nitrogen content, P: leaf phosphorus content, Area: leaf area, Fresh mass: leaf fresh mass, SLA: Specific leaf area,
 544 Thickness: leaf thickness, DE: deciduousness, H_{max} : maximum height, WD: wood density, Seed mass: weight of the seed.
 545

546 Overall, we find that 1) trait-environment relationships are similar for most traits across
 547 lowland and montane tropical American forests; 2) lowland forests show significant and larger
 548 changes in more community traits than montane forests; 3) across the forests and for the full
 549 community and survivor assemblages, the abundance of deciduous species is increasing, with
 550 accompanying increases in leaf photosynthetic capacity and decreases in leaf area and leaf
 551 thickness. However, the recruiting communities in the lowland forests have on average decreased
 552 in the abundance of deciduous species, leaf carbon and nitrogen content and wood density; and 4)
 553 crucially, for the full tree community and survivors assemblages most of these traits are changing at
 554 only a fraction of the rate required to maintain equilibrium with climate. Notably, the recruiting
 555 communities show the best tracking of a changing climate.

556 The community trait shifts were similar for the full community and the survivor assemblages
 557 but generally differed from those of the recruitment and mortality assemblages. This is likely because
 558 the trait shift responses of the full community and survivor assemblages are dominated by large
 559 individuals that continued growing throughout the study period. Although large individuals would
 560 have the highest impact on ecosystem functioning (45) due to their size and long lifespans, the
 561 functional responses of the community to ongoing climate change may be more evident when
 562 analysing the functional traits of recruiting individuals. Moreover, it is the recruiting communities that
 563 will eventually dictate future community traits. There is a mismatch in trait responses to climate
 564 change between the recruiting assemblage and both the full community and survivor assemblages.
 565 This mismatch is most pronounced with respect to the abundance of deciduous species, leaf carbon,
 566 and maximum adult height. The recruiting assemblages select for lower leaf carbon, lower
 567 deciduousness, and species with shorter maximum heights. This finding suggests a decoupling in
 568 trait space between the functional trait characteristics of old forests we see in the present, and the
 569 possible future functional composition of tropical American forests. The selection for low leaf nitrogen
 570 in the mortality and recruitment assemblages raises the question of whether and to what extent such
 571 recruitment assemblages with low leaf nitrogen content will be able to survive to larger adult sizes
 572 (e.g. 55, 56), especially across montane forests where there is a stronger mismatch (Fig. 3 and Fig.
 573 4). Such a decoupling in trait space between the survivor and recruitment assemblages could
 574 potentially indicate the slow beginnings of forest level adaptation to new climatic conditions, which
 575 is likely to impact the functioning of tropical forest ecosystems (57).

576 The full community, survivor and recruiting assemblages often show more changes in traits
 577 in lowland than montane forest. Lowland forests are highly dynamic and harbour a high functional
 578 trait diversity that potentially allows for selection from a wider pool of trait values under climate stress.

579 There has been a larger increase in atmospheric VPD in lowland forests than in montane forests,
580 caused by more pronounced increases in temperature over the last 40 years, which could partially
581 explain the shift of a larger number of community functional traits in lowland than montane forests
582 (58). Larger increases in VPD and more severe droughts appear to have modified the community
583 composition of lowland forests more strongly than that of montane forest, towards a set of species
584 better adapted to drier and hotter conditions, which could be due to the mortality of more vulnerable
585 species (46). Recent work across sites in the Amazon and Andes also suggest an important impact
586 of increasing temperatures and declines in water availability on changes in tree trait composition
587 (59). A shift in community composition towards more deciduous species in both lowland and
588 montane forests appears to have caused increases in photosynthetic capacity, leaf chemistry traits
589 such as leaf carbon and phosphorous content, and species maximum height through time.

590 The abundance of deciduous tree species has generally increased in tropical American
591 forests over the last 40 years (Fig. 2C). For the full-community and recruiting assemblages this was
592 accompanied with a general decline in leaf thickness. Temporal increases in VPD have favoured
593 increases in the proportion of deciduous species, especially across montane forests, and increases
594 in MCWD have promoted decreases in leaf thickness. Overall, deciduous species tend to have
595 acquisitive leaf traits with higher leaf nitrogen and phosphorus, photosynthetic capacity and
596 photosynthetic nitrogen-use efficiency, especially under hydric stress⁵⁶, than evergreen species (60,
597 61). The pattern observed across tropical American forests could be attributable to leguminous
598 nitrogen fixing species that dominate in dry forests which are often deciduous and with higher
599 photosynthetic nitrogen-use efficiency (62). This is consistent with a previous report for West African
600 tropical forests, where increasing drought stress co-occurred with an increased abundance of
601 deciduous species, and where changes in deciduousness explained changes in other morphological,
602 structural and leaf chemistry traits (50). However, the distribution of tree species with such trait
603 syndromes may be limited by the spatial heterogeneity in essential soil nutrients, such as nitrogen
604 and phosphorus (63). Thus, increase in deciduousness is expected to be one adaptation strategy,
605 especially in dry tropical forests with more seasonal precipitation regimes than wetter tropical forests.
606 Yet, increased deciduousness may be limited by soil fertility in some areas, such as in infertile south-
607 eastern Amazonia, where short-lived deciduous leaf construction is a too-costly strategy.

608 With increasing temperatures and reduced water availability, we expected an increase in
609 abundance of deciduous species to also be reflected in the recruitment assemblage (50). However,
610 the decline in abundance of deciduous species in the recruits assemblage indicates potential shifts
611 in phenological strategies towards more conservative strategies in response to increasing
612 temperatures or altered precipitation patterns. However, we did not find a significant selection
613 against deciduous species in the mortality assemblage. This suggests that a combination of drought
614 avoidance and drought resistance strategies (34) could both be playing an important role as means
615 of adaptation to a warming climate across lowland and montane tropical forests. Other factors may
616 be promoting the observed change in community-mean traits, such as defaunation, which could be
617 an important driver of the observed changes in dispersal traits across time (64). Some wetter regions
618 (e.g., central Amazonia) show increases in seed mass for the full community, with the mortality
619 assemblage showing significant declines in individuals with smaller seeds in the lowlands (Fig. 4C).
620 However, drier regions (e.g., southern and eastern fringes of Amazonia) and montane forests show
621 a slight predicted decline in seed mass (Fig. S3 D). These changes may be an indicator of strong
622 defaunation pressure (65) as spatial predictions of decreases in seed mass broadly match spatial
623 patterns of high defaunation (66), especially in those more accessible areas of Mesoamerica, and
624 both south and eastern Brazil. They could also be driven by climatic factors as the observed changes
625 are consistent with a shift from endozoochory (animal dispersal) to anemochory (wind dispersal),
626 with the latter exhibiting smaller seeds than those dispersed by animals and being more prevalent
627 in drier biomes (39).

628 We investigated the impact of macroclimate on the changes in functional trait composition of
629 tropical forests. However, such macroclimate conditions may not directly mirror the microclimatic
630 conditions found under the forest canopy (67). This is of particular importance when investigating
631 the effects of a changing climate, especially on the forest's recruiting assemblages, which tend to
632 occupy the space below the canopies of the older larger trees. Ultimately, such microclimatic

633 conditions may play an important role for determining the responses of below canopy plants to a
634 changing climate (68, 69, 70) and therefore on the rate of change in community trait composition of
635 the recruiting assemblages. Hence, microclimatic conditions at the plot level may partly explain the
636 differences in trait shifts between the full community and survivors assemblages and the recruiting
637 assemblages.

638 Changes in single traits do not always behave as expected from theory. For instance, it would
639 mechanistically be expected that increasing drought would cause plant communities to shift to
640 species with higher wood density and thicker leaves or that the abundance of deciduous species
641 increases across time. Such coordinated changes may not readily happen in the community as it is
642 whole phenotypes that are changing, i.e. particular combinations of traits, rather than isolated traits.
643 Moreover, coordination of different strategies could allow for alternative adaptations to the same
644 drivers. For example, drier conditions might encourage deciduousness combined with low wood
645 density and thin leaves (drought avoidance), or evergreenness combined with high wood density
646 and thicker leaves (drought tolerance). The favoured combination(s) may depend on forest
647 seasonality patterns and soil nutrients. Furthermore, not all trait combinations may be present in any
648 given regional species pool, even in species-rich biomes, which may limit the shifts in community
649 traits that can occur at any given time as a response to environmental change. Other factors may
650 also contribute to trait shifts or a lack thereof across forest communities, such as soil conditions (71),
651 biotic interactions (e.g., animal-plant interactions) (72) and wind disturbance (73). Our analyses
652 represent community-wide responses mainly based on trait information at the species and genus
653 level; traits may also express intraspecific plasticity that we are unable to assess here given the
654 scale and multidecadal nature of the study. Some traits may show more or less plasticity than others
655 and species intraspecific variation may contribute to adaptation to a changing climate (74, 75).

656 In conclusion, we find that overall community composition changes are leading to small shifts,
657 ~10% of the expectation given climate change, in the overall community trait composition. These
658 shifts are primarily driven by variation in growth rates of existing trees, rather than by new recruitment
659 or mortality. However, we observed larger changes for the recruiting assemblage, directionally
660 tracking climate at an average of 21%, which can potentially contribute to keeping these forests
661 closer to, although still far from the equilibrium with climate. Trees are long-lived organisms with slow
662 turnover rates compared to the rate of climate change and this partly explains the differences
663 observed in community trait shifts between the full community and those of the recruiting
664 assemblages. There are specific areas where there seems to be a larger lag in forest responses to
665 climate changes, especially in the Maya forest in Mesoamerica (76), and both the Atlantic forest and
666 the southern Amazon forest in Brazil (77), which have become increasingly fragmented over time.
667 Consequently, across these regions, edge effects and in general a more constrained physical
668 environment may be impacting the capacity of forests to adjust to new climate conditions (28, 78).
669 Our analysis demonstrates that tree community composition is shifting to track climate change, but
670 that the overwhelming onus would have to be on within-species variability and trait plasticity (79, 80)
671 to adequately track climate change. However, the changes in climate are likely to be too fast for
672 adaptive phenotypic plasticity to keep track, especially in environments with low climatic
673 heterogeneity (79, 80). Hence it is overwhelmingly likely that tree species composition and functional
674 properties of tropical American forests (and probably all tropical forests) are increasingly out of
675 equilibrium with local climate. Such disequilibrium almost certainly increases vulnerability.

676 **Materials and Methods**

677 **Plot data**

678 Our study focuses primarily on tropical American forests, with some overflow into adjoining
679 subtropical and warm temperate regions of Brazil and Mexico. We gathered tree-by-tree vegetation
680 census data for 254,307 individual trees from 415 vegetation old-growth forest plots across 11
681 countries, spanning a wide range of environmental conditions and elevations from sea level to >3000
682 m elevation, with at least two censuses recorded (on average 5.7 censuses per plot) between the
683 years of 1980 and 2021 (Table S1). The modal plot size was 1 ha (mean 0.88 ha); all plots were
684 located in structurally intact forests with no signs of direct anthropogenic impacts. In each, all woody
685 plants with a diameter ≥ 10 cm at breast height (DBH) or above buttress roots were measured.

686 Overall, identification rates of individuals exceeded 85% to the species level and 99% to the genus
687 level (Fig. S8). Data were obtained through the ForestPlots network (www.ForestPlots.net) (83) and
688 the MONAFOR Mexican plots network (forestaes.ujed.mx/monafor). We classified vegetation plots
689 as either lowland or montane forests. We based our classification on literature that uses 500-800
690 metres above sea level (masl) as the threshold between lowland and montane climatic conditions
691 (29, 83). In our dataset there were two clear elevational groups, where plots <604 masl (261 plots)
692 were classified as lowland forests, and plots >730 masl (154 plots) were classified as montane
693 forests.

694 695 **Trait data**

696 Tree functional trait data were obtained for several plots from local field collections carried out by
697 collaborators from where plots are located (84, 85, 86), the Global Ecosystems Monitoring network
698 (GEM; gem.tropicalforests.ox.ac.uk) (87), and the ForestPlots network (www.ForestPlots.net) (82)
699 in addition to databases from BIEN (bien.nceas.ucsb.edu), TRY (www.try-db.org) (88) and Diaz et
700 al. (89). The plant traits are related to fundamental aspects of leaf chemistry, photosynthetic
701 capacity, leaf morphology, plant and organ size, and phenology (Table S2). We gave priority to trait
702 data from the GEM and ForestPlots networks; when these were not available, we used the
703 databases from TRY (www.try-db.org), Diaz et al. (89), and the BIEN (bien.nceas.ucsb.edu)
704 network's database. When more than one trait value per species was available, we used the trait
705 mean at the species level for subsequent analysis. We aimed to cover at least 70% of the basal area
706 of each plot with trait data at species or genus level, often covering more than that (Fig. S9). When
707 species-level trait data were unavailable, we used the mean genus-level data. Hence, our analysis
708 could be seen as more representative of the genus-level trait responses. When achieving at least
709 70% coverage was not possible for a given trait, such a plot was left out of the analysis for the
710 specific trait. All species names were standardised following the Taxonomic Name Resolution
711 Service (TNRS; <https://tnrs.biendata.org>).

712 713 **Climate data**

714 We investigated the role that long-term climate and its changes play on determining the trait
715 community composition across tropical American forests by gathering climatic data on the mean
716 annual values of temperature (T_{mean}), maximum climatic water deficit (MCWD) (47), vapour pressure
717 deficit (VPD_{mean}) (46), standardised precipitation-evapotranspiration index for a 12-month window
718 ($SPEI_{12}$) (48) and dry season length. We calculated the long-term climate conditions as the mean
719 annual values for the metrics described above between the years 1980 to 2021 (i.e., full time period).
720 All climatic variables were derived from the TerraClimate dataset (31) and had an original spatial
721 resolution of $\sim 4 \times 4$ km at the Equator. We selected the TerraClimate product because of its accuracy
722 and coverage. TerraClimate has higher spatial resolution and greater accuracy when compared to
723 other historical climatic datasets such as Climatic Research Unit data (44). The dry season length
724 was calculated as the average annual number of consecutive months with rainfall below 100 mm
725 (88). The MCWD was included as it is a metric for drought intensity and severity that has been shown
726 to impact vegetation growth and survival (50). MCWD is thus defined as the most negative value of
727 the climatological water deficit (CWD) each year. We converted the MCWD so that positive values
728 indicate increases in water stress. Equally, the SPEI reflects drought severity, but its multi-scalar
729 nature enables the identification of different drought types and severity (91). VPD is an indicator of
730 atmospheric aridity, acts as a key environmental driver of plant transpiration, and reduces plant water
731 use efficiency (92). We then tested the correlation between all pairs of climatic variables (full-term
732 and their changes); all had Pearson's correlation coefficients $|\leq 0.70|$ apart from dry season length
733 which was highly correlated to MCWD (Fig. S1), and we thus dropped dry season length and its
734 change to avoid confounding model coefficients (93). We also obtained the same variables outlined
735 above for a period covering 1980-2005 (see models below). We calculated the yearly rate of change
736 of the climatic variables ($\Delta r T$, $\Delta r MCWD$, $\Delta r VPD$, and $\Delta r SPEI_{12}$) to standardise for a different time
737 between censuses for different plots and avoid the bias due to inter-annual short-term variability that
738 occurs in addition to the long-term change. To this end, we fitted a linear model predicting the climate
739 variable value as a function of time (year) and used the slope as the predicted annual rate of change
740 (Δr).

741

742 Soil variables are relevant predictors of vegetation distribution and are related to functional
743 trait composition (14). Variation in soil properties could modify the rate of change in response to
744 environmental change (94). Hence, in our models, we included soil clay (%), sand (%), cation
745 exchange capacity (CEC; mmol) and pH information averaged from the first 30 cm of the soil surface.
746 The soil data was derived from the SoilGrids global dataset (www.soilgrids.org).

747 All climatic and soil variables were numerically centred around the mean value before model
748 fitting. The study area that was used to extract the climatic and soil data and to make spatial
749 predictions was delineated using the European Space Agency Land Cover CCI Product (81) using
750 all land use classes defined as tree or shrub cover classification.

751 **Trait CWM calculation**

752 The most dominant species are expected to drive ecosystem processes using their traits as
753 described by the mass ratio hypothesis (45). Therefore, for each of the plant functional traits t and
754 plots p per census time we calculated their community-level weighted mean (CWM) using the
755 species basal area as the weighting factor: $CWM_{xp} = \sum_{i=1}^S BA_{ip} \times x_i$. Here BA_{ip} is the basal area
756 of species i in plot p , with x_i representing the average trait value of species i . Before calculating the
757 trait CWM, when more than one value per trait was available for a species we averaged the trait
758 value at the species level; when the species had no trait values, we averaged the trait values at the
759 genus level. Although species show some degree of intraspecific trait variation, research suggests
760 it is relatively small compared to the trait variation found across forest tree species (95). Moreover,
761 given the vast majority of functional trait data has only been collected in the last decade, it is not yet
762 possible to evaluate the magnitude of intraspecific trait shifts across the spatial extent of the study area.
763 The trait CWM is an indicator of mean canopy properties as basal area and crown area are highly
764 related, and the latter indicates the amount of canopy area occupied by a specific trait (96). In the
765 case of phenological strategy, as it was obtained as a categorical variable (deciduous or not
766 deciduous), we calculated the percentage of basal area that is deciduous.

768 **Understanding trait CWM-Climate relationships and the effects of climate change for driving 769 trait CWM changes**

770 All statistical models described below are linear models fitted under a Bayesian modelling framework
771 with the 'rstanarm' R (97) package. All models were run with normal diffuse priors with a mean of 0
772 and 2.5 standard deviations (sd) to adjust the scale of coefficients and 10 sd to adjust the scale of
773 the intercept. The models were run with three chains and 3,000 iterations. We computed the HDI
774 (highest density interval), resulting in the range containing the 95% most probable effect values (98,
775 99), and considered a result significant when the HDI did not overlap 0. The CWM change models
776 (a separated model per trait) used a Gaussian distribution, weighting the observations by the number
777 of censuses, census time and plot size (9).

778 To understand the current trait-climate relationships across forests of the tropical Americas, for each
779 plant trait we modelled the trait CWM as a function of the T_{mean} , VPD_{mean} , $MCWD_{\text{mean}}$, the $SPEI_{12}$
780 (from the full time period data) and the soil covariates, with each one of the climatic variables
781 interacting with forest type (lowland or montane) (here onwards referred to these models as M1).
782 Because the studied functional trait values are always positive and often have a long-tailed
783 distribution, the current Trait-Climate relationship statistical models used a Gamma distribution and
784 log link function, using a weighting given the number of censuses, plot size and census time. We
785 tested the model predictions using a spatial K-fold cross-validation approach where the spatial
786 groups were determined based on the group ID described below. All models were developed in the
787 R language for statistical computing (97).

788 We next analysed the climatic drivers of shifts in each functional trait given observed changes
789 in climate over the past 40 years for the full community and survivors assemblages (415 vegetation
790 plots), for the recruiting community (377 plots) and mortality community (375 plots). We define the
791 recruiting community as individuals that passed the threshold of 10 cm DBH in a given census and
792 were not present in the past census for a given plot. The mortality community is defined as those
793 individuals of a plot who were alive in a previous census but dead in the following census. We

794 calculated the temporal changes in trait CWM at the plot level as the annual rate of change (Δr of
795 the trait CWM) to standardise for a different time between censuses for different plots. To this end,
796 we fitted a linear model predicting the trait value as a function of time (year) and used the slope as
797 the predicted annual rate of change (Δr). We then modelled the Δr CWM trait as a function of Δr
798 of the climatic variables described above, each one of these interacting with forest type and also
799 included the soil characteristics to account for their possible effect on the observed CWM trait
800 changes (hereafter referred to these models as M2). We weighted the observations by the number
801 of censuses, census time and plot size. We used the M2 models to predict and spatially project the
802 changes in trait composition across Latin American forests over the past half-century.

803 We tested for spatial autocorrelation effects in the CWM traits using the Moran's I test and
804 found significant effects (P -value <0.05). Therefore, for each trait CWM we calculated the spatial
805 distance at which the spatial effect decreased, finding that a distance of nine kilometres for the full
806 community and survivors assemblages (1 km for recruitment and mortality) captured most of the
807 spatial autocorrelation effect. Based on this we generated an identification (ID) for each group of
808 plots that were at maximum nine kilometres away (or 1 km respectively) from each other and
809 included such group ID as a random factor (14) in the statistical models to account for the spatial
810 autocorrelation and visually checked the model residuals for any obvious patterns (5).

811 812 **Understanding shifts in trait CWM**

813 We used the annual rate of change (Δr) of the trait CWM of the full, survivors, recruiting and mortality
814 assemblages to investigate if the rate of trait changes for the overall forests (lowland and montane
815 together), for the lowland forests alone and the montane forest alone, was significantly different from
816 0. We did the same to understand if there were important differences between the rate of change
817 between lowland and montane forests. To this end we carried out a Bayesian version of a typical T-
818 test analysis using Bayesian estimation (99, 100). The Bayesian estimation was done using the
819 'BEST' package for R, with normal priors with mean for μ (the mean of rate of change) of 0 and a
820 standard deviation for μ of 10. We used broad uniform priors for σ (standard deviation), and a
821 shifted-exponential prior for the parameter ν which describes the normality of the data within the
822 group. As above, here we calculated the HDI containing the 95% most probable effect values and
823 considered a result significant when the HDI did not overlap 0.

824 **Understanding if forest community traits are tracking climate changes.**

825 The process outlined below was carried out for the full community, the survivors and recruiting
826 assemblages only as the mortality ones are not tracking climate. We first built the same type of
827 statistical models constructed above for the current trait-climate relationships (M1) but using only
828 plot and climatic data from between 1980 and 2005, including also the soil variables described above
829 (from now on called M1.1). This was done as there is potential for entanglement when using the
830 same data for spatial and temporal analysis. The plot data up to year 2005 were still representative
831 of the plot distributions across tropical Americas and helped separate the temporal effect from the
832 spatial relationships.

833
834 We used the M1.1 Trait-Environment statistical models and obtained predictions of the trait
835 CWM to a new set of climatic conditions composed of the 1980-2005 climate plus the observed
836 climate yearly rate of change across the study period (here onwards M2). We then calculated the
837 difference between the trait CWM obtained with the M1.1 and M2 models to obtain the expected trait
838 CWM change. Lastly we compared the expected trait CWM calculated above with the observed Δr
839 CWM trait (i.e., from section 'Understanding trait CWM-Climate relationships and the effects of
840 climate change for driving trait CWM changes'). This allowed us to understand the expected shift in
841 mean trait values given the 1980-2005 trait-climate relationship in comparison to the observed trait
842 changes across time (i.e., from 1980-2021). Thus, the 1980-2005 trait-climate relationship shows
843 how much the tree communities would need to change to keep up with climate changes, and the
844 observed trait changes across time show how much they have actually changed. We compared the
845 observed and expected average change trait CWM across the sampling vegetation plots to
846 understand the magnitude of their difference for lowland and montane forests and calculated their

847 change ratio. To understand if there was a significant difference between observed and expected
848 community trait changes, we carried out a Bayesian version of a typical T-test analysis using
849 Bayesian estimation as described above and considered significance when at least 95% of the HDI
850 values were greater than zero (99, 100).

851 We also created map predictions of the 1980-2005 M1.1 trait-climate model across tropical
852 American forests by predicting this model to a climate change scenario that was composed of the
853 observed climate (1980-2005) plus the yearly rate change observed. We then subtracted the original
854 map predictions (those made with the M1.1 models without changes in climate conditions) to obtain
855 the expected CWM trait changes at the pixel level (in the map) for across forests in tropical America.
856 Then we calculated the ratio of the observed, i.e., spatial predictions of the trait changes observed
857 across time (from M2 models), versus expected and converted to percentage change relative to the
858 1980-2005 condition to understand if and to what extent the observed trait changes are tracking
859 (values above zero) or not (values of zero) the expected changes given the observed changes in
860 climate or shifting in opposite direction than expected (values below zero). All statistical analyses
861 were carried out in R.

862 **Creating the spatial predictions (maps)**

863 All maps were generated by predicting the focus model (i.e., M1, M1.1, M2) to the study area. The
864 study area was delineated using the European Space Agency Land Cover CCI Product (81) using
865 all classes having tree cover classification and numbered from class 50 to class 100 as suggested
866 here: <http://maps.elie.ucl.ac.be/CCI/viewer/index.php>. To avoid extreme values in the maps, given
867 some extreme climate and soil values inherent to the climate and soil data, we truncated the map
868 predictions to contain the 90 percentile predicted value as the maximum instead of the 100% which
869 allowed us to eliminate the outlier values. The maps were created in the R platform.

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1201

1202 **Data availability**

1203 The vegetation census and plant functional traits data that support the findings of this study are
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1208

1209 **Supplementary Materials**

1210 Supplementary Text

1211 Figs. S1 to S7

1212 Tables S1 to S9